Multi-century Dynamics of the Climate and Carbon Cycle under Both High and Net Negative Emissions Scenarios

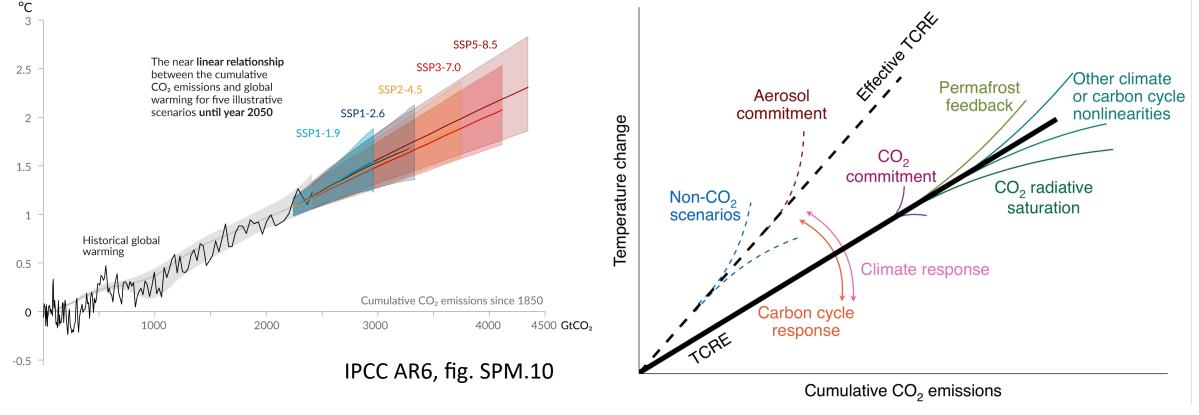
Charles D. Koven¹, Vivek K. Arora², Patricia Cadule³, Rosie A. Fisher^{4,5,6}, Chris D. Jones⁷, David M. Lawrence⁴, Jared Lewis⁸, Keith Lindsay⁴, Sabine Mathesius⁹, Malte Meinshausen^{8,10}, Michael Mills⁴, Zebedee Nicholls^{8,10}, Benjamin M. Sanderson^{4,5}, Roland Séférian¹¹, Neil C. Swart², William R. Wieder^{4,12}, Kirsten Zickfeld⁹

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 ¹Climate and Ecosystem Sciences Division, Lawrence Berkeley National Laboratory, Berkeley, California, USA
²Canadian Centre for Climate Modelling and Analysis, Environment and Climate Change Canada, University of Victoria, Victoria, British Columbia, Canada ³IPSL, CNRS, Sorbonne Université, Paris, France
⁴Climate and Global Dynamics Laboratory, National Center for Atmospheric Research, Boulder, Colorado, USA
⁵Centre Européen de Recherche et de Formation Avancée en Calcul Scientifique, (CERFACS), Toulouse, France
⁶Évolution & Diversité Biologique, University of Toulouse Paul Sabatier III, Toulouse, France.
⁷Met Office Hadley Centre, Exeter, UK
⁸Climate and Energy College, School of Geography, Earth and Atmospheric Sciences, The University of Melbourne, Parkville, Victoria, Australia
⁹Department of Geography, Simon Fraser University, Burnaby, British Columbia, Canada
¹⁰Climate Resource, Victoria, Australia
¹¹CNRM (Université de Toulouse, Météo-France, CNRS), Toulouse, France
¹²Institute of Arctic and Alpine Research, University of Colorado, Boulder, Colorado, USA

Proportionality of warming to cumulative CO2: does it have limits, and if so, what are they?

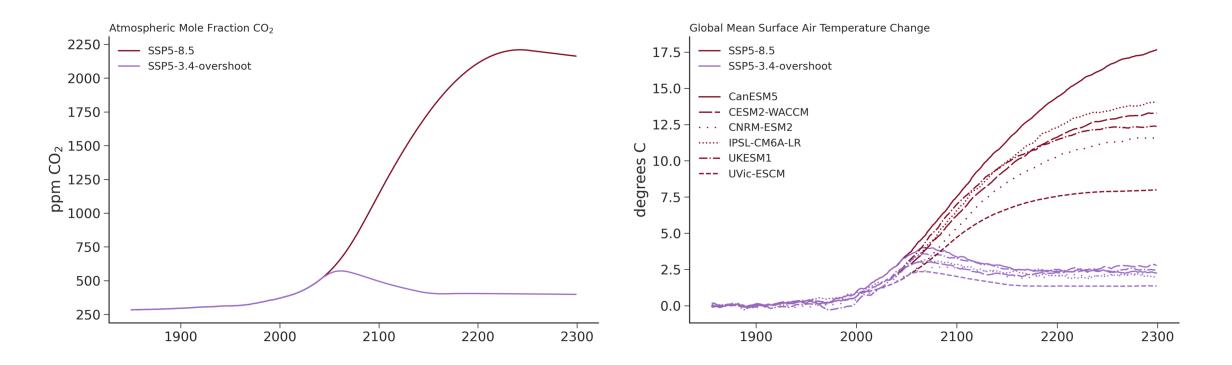
Global surface temperature increase since 1850-1900 (°C) as a function of cumulative CO₂ emissions (GtCO₂)



Matthews et al., 2020

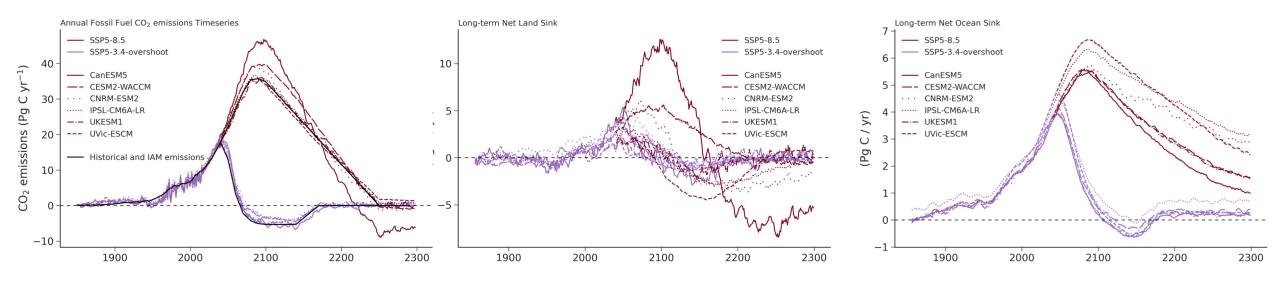
Perhaps most importantly, does the relationship hold under net-negative CO₂ emissions?

Comparing two long-term scenarios: veryhigh emissions, and peak-and-overshoot

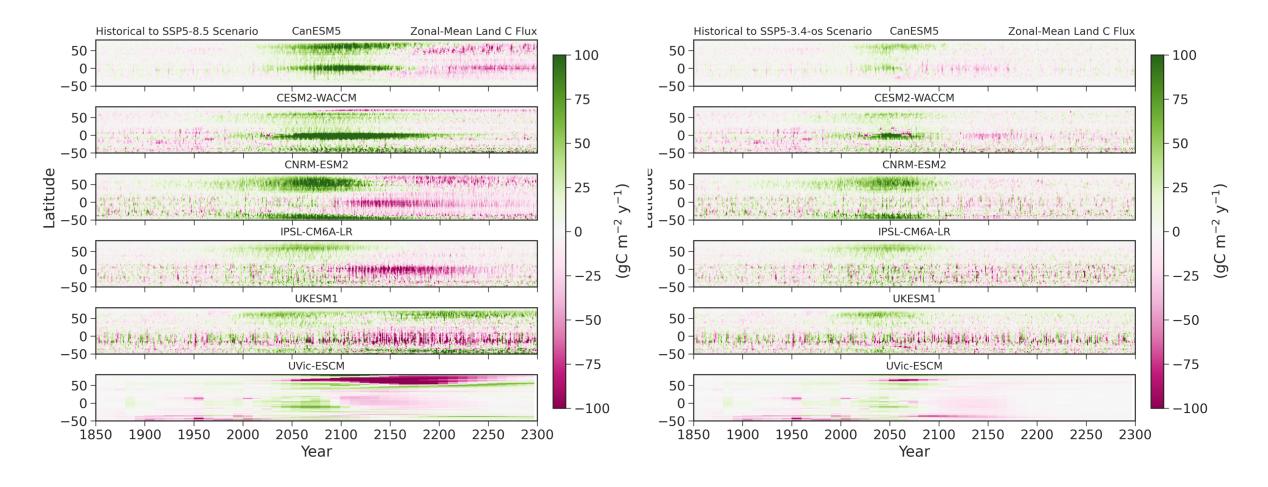


All unlabeled figures in talk are in revisions manuscript in review at https://doi.org/10.5194/esd-2021-23

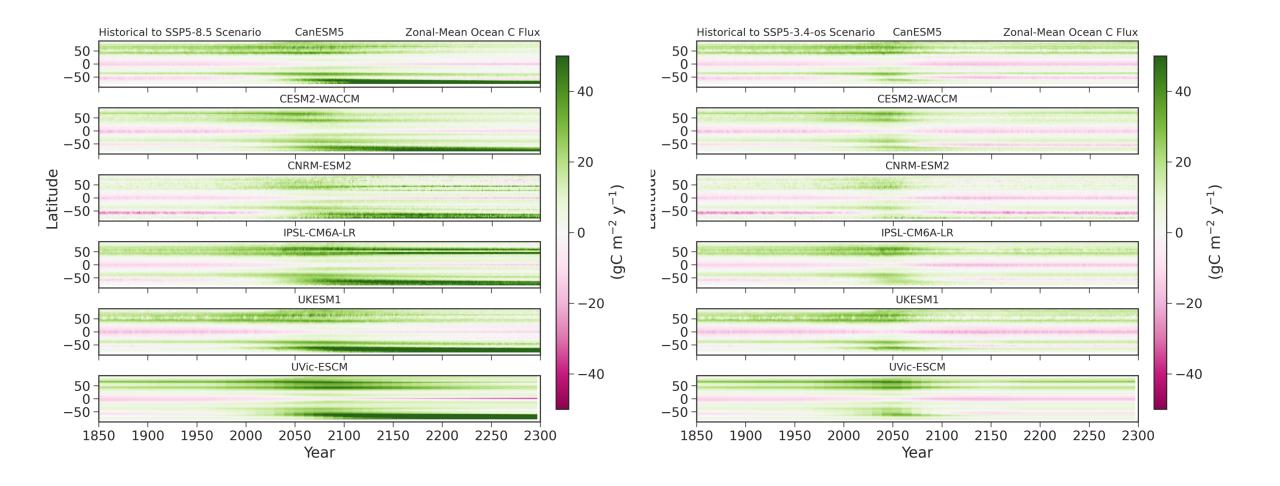
Land and ocean carbon fluxes: land switches from sink to source under both scenarios, ocean follows emissions and switches to source during overshoot period



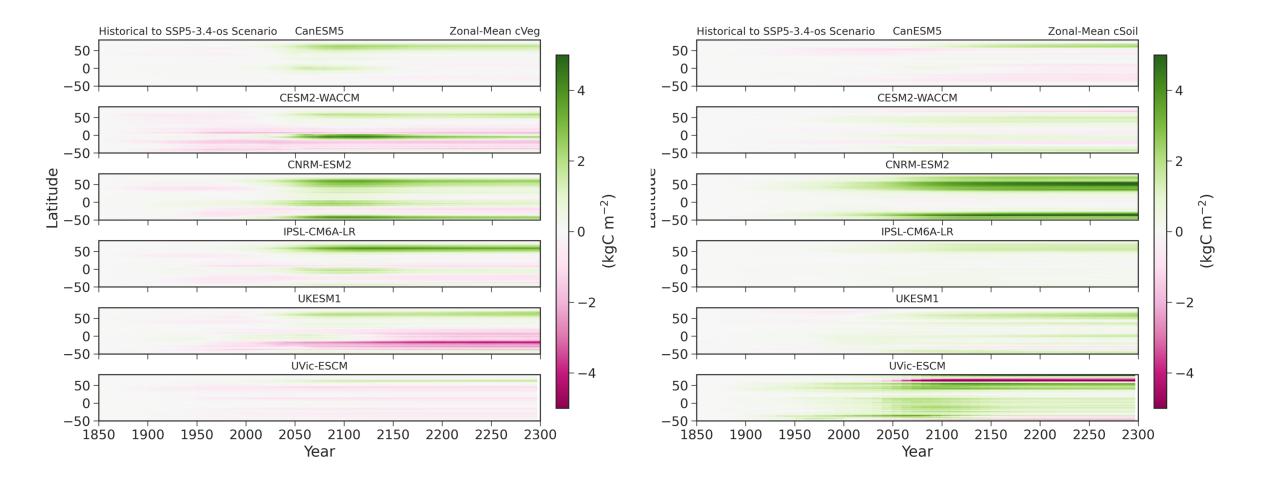
Looking regionally, models disagree on the timing, location, and strength of feedbacks over land, particularly under high very-high emissions scenario.



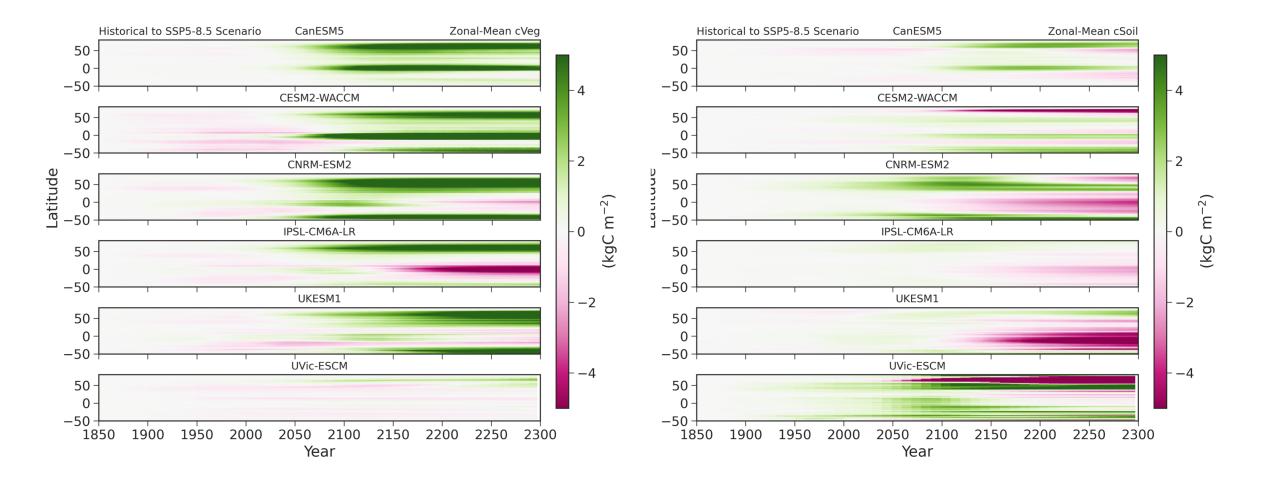
For ocean, models agree on timing and location, only some disagreement on magnitude of feedbacks.



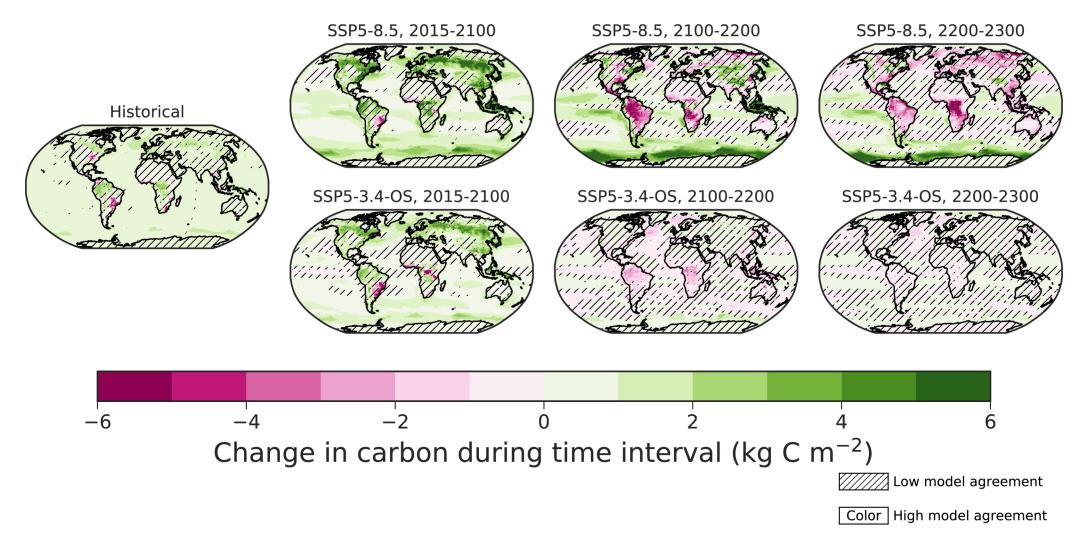
Comparing the veg and soil carbon responses, again wide disagreement on timing, location, magnitude and veg/soil split, but also whether feedbacks are reversible after overshoot.



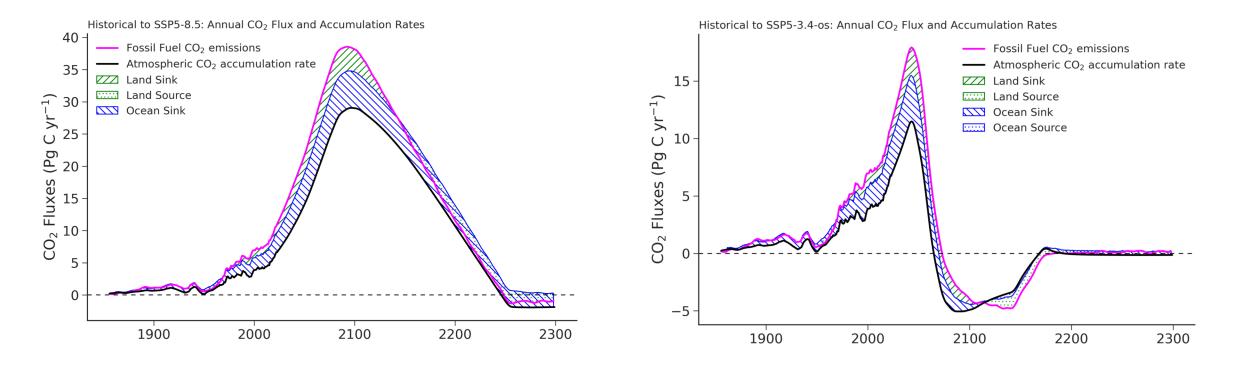
Comparing the veg and soil carbon responses under very high emissions, all models agree that high latitudes gain veg. carbon, but disagree on much else.



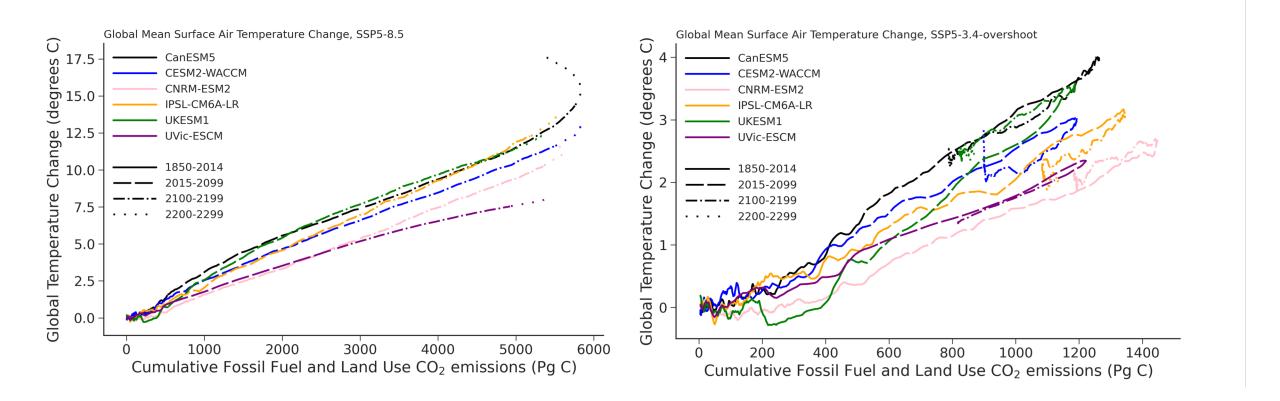
Maps of integrated carbon change show that there are also key longitudinal differences, e.g. that Amazon tends to show earlier vulnerability than African and Maritime tropical forests.



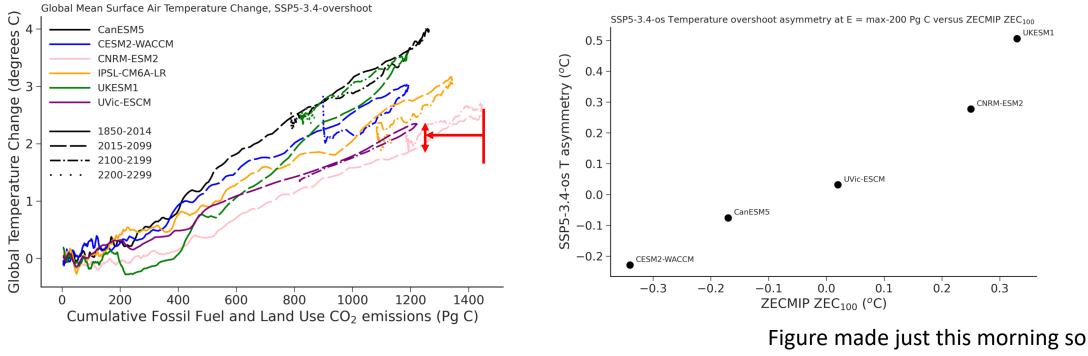
Plotting all fluxes into a single plot allows seeing the relative lags between carbon emissions and feedbacks.



Proportionality of warming to cumulative emissions roughly holds, under both scenarios, with some exceptions.



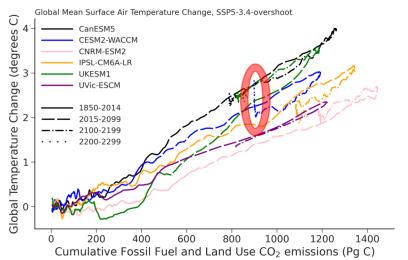
Overshoot asymmetry consistent with 100-year Zero Emissions Commitment (ZEC_{100})



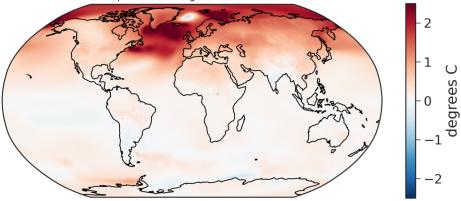
not in revised manuscript

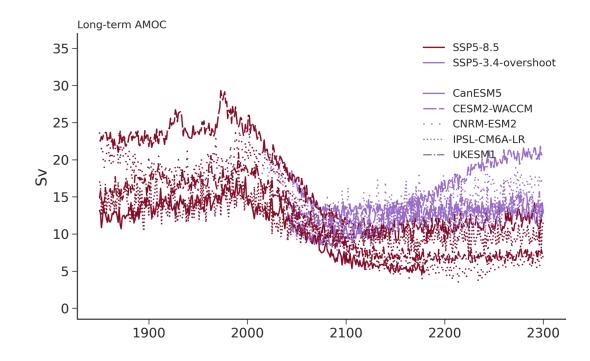
ZEC₁₀₀ values from MacDougall et al., 2020 (https://doi.org/10.5194/bg-17-2987-2020); IPSL-CM6A-LR not included in ZECMIP

23rd-century warming in CESM2 under SSP5-3.4overshoot concentrated in northern hemisphere, consistent with driving by AMOC recovery

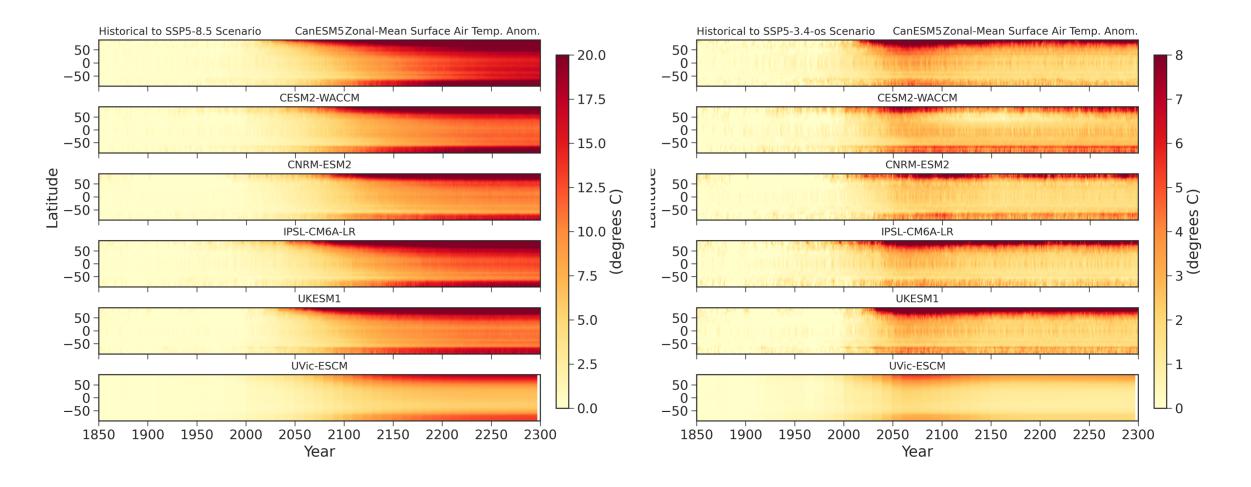


CESM2-WACCM Temperature change from the 22nd to 23rd centuries





Zonal-mean temperature changes show expected patterns of polar amplification, and also shows signature of 22nd-century northern mid-high latitude cooling in CESM2



Conclusions

- Carbon cycle responses in both long-term scenarios include reversals of global sink to source in land and/or ocean, for very different reasons
 - Land may become source under both net-negative and long-term very-high emissions, whereas ocean only become transient source under strong net-negative emissions
- Wide model disagreements in land sink are particularly strong under high emissions scenarios, model disagree on timing, location, strength, and veg/soil partitioning. Somewhat better agreement in overshoot scenario
- Ocean models agree much more on patterns, but ensemble spread still diverges after 2100
- Overall the proportionality of warming to cumulative emissions still generally holds under both scenarios
- Asymmetry under overshoot largely consistent with ZEC
- Long-term divergence from proportionality to warming can be seen in both scenarios though, e.g. due to AMOC recovery in CESM2 under overshoot, or lagged warming in CanESM5 under very-high-emissions